

Step by step to PaRIS – Airborne Carrier Phase Altimetry using reflected GPS signals

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GPS Sea Ice - Introduction

- Geometry
- Altimetry

Tropospheric Bias

- Model
- Observation
- Results

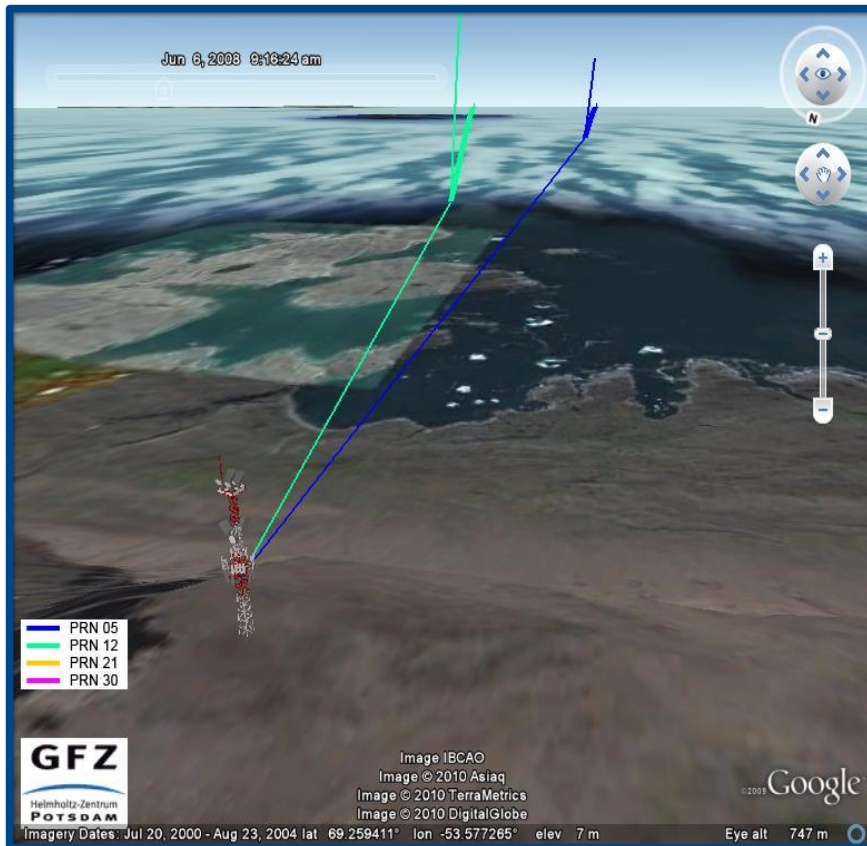
Ionospheric Bias

- Model
- Observation
- Results

Summary & Outlook

GPS Sea Ice - Introduction

Geometry



Receiver

- coast high above the ocean
- Earth fixed position

Transmitter

- at low elevation
- rising and setting

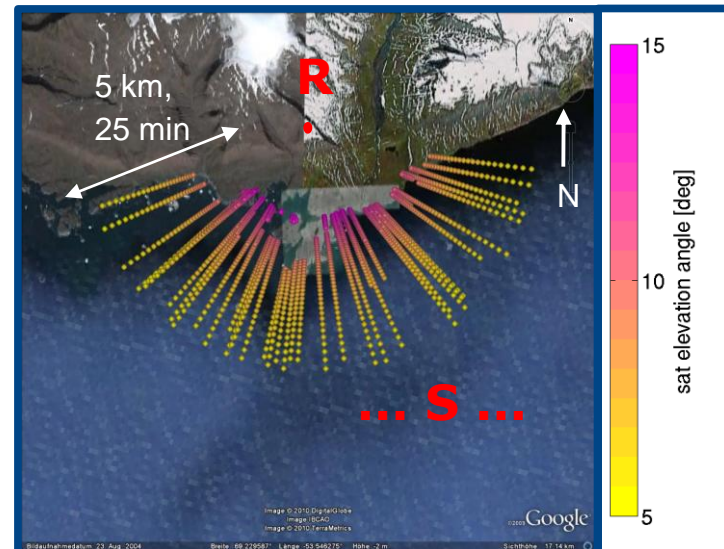
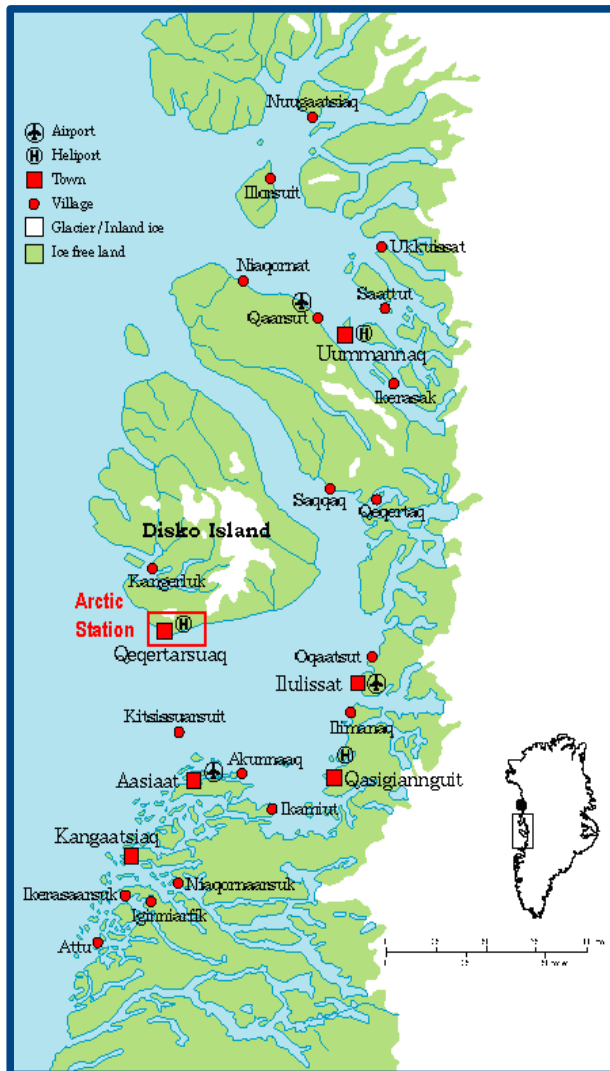
Reflection Track

- specular points
- on the ocean surface

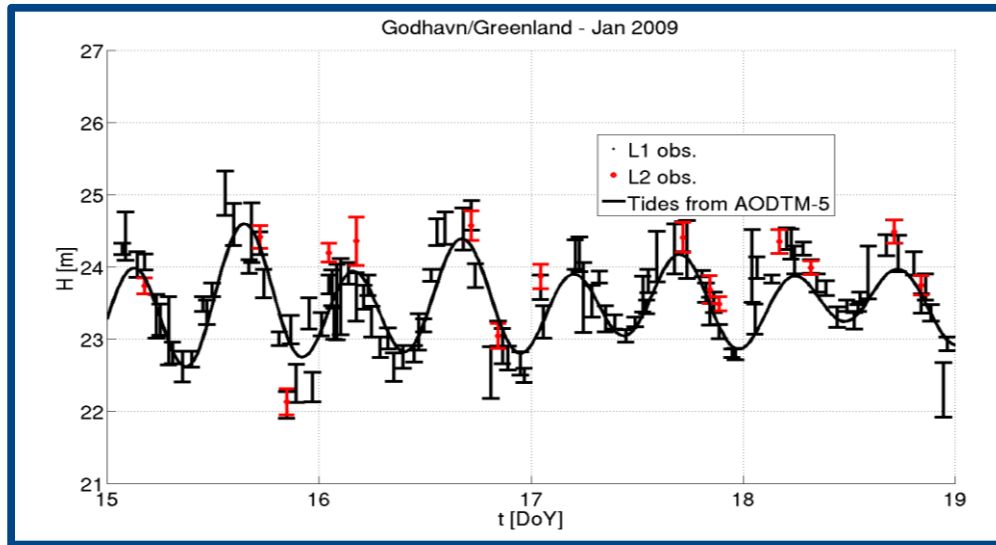
Geometry

Reflection Tracks

- fan-shaped area
- low satellite elevation angle



Altimetry

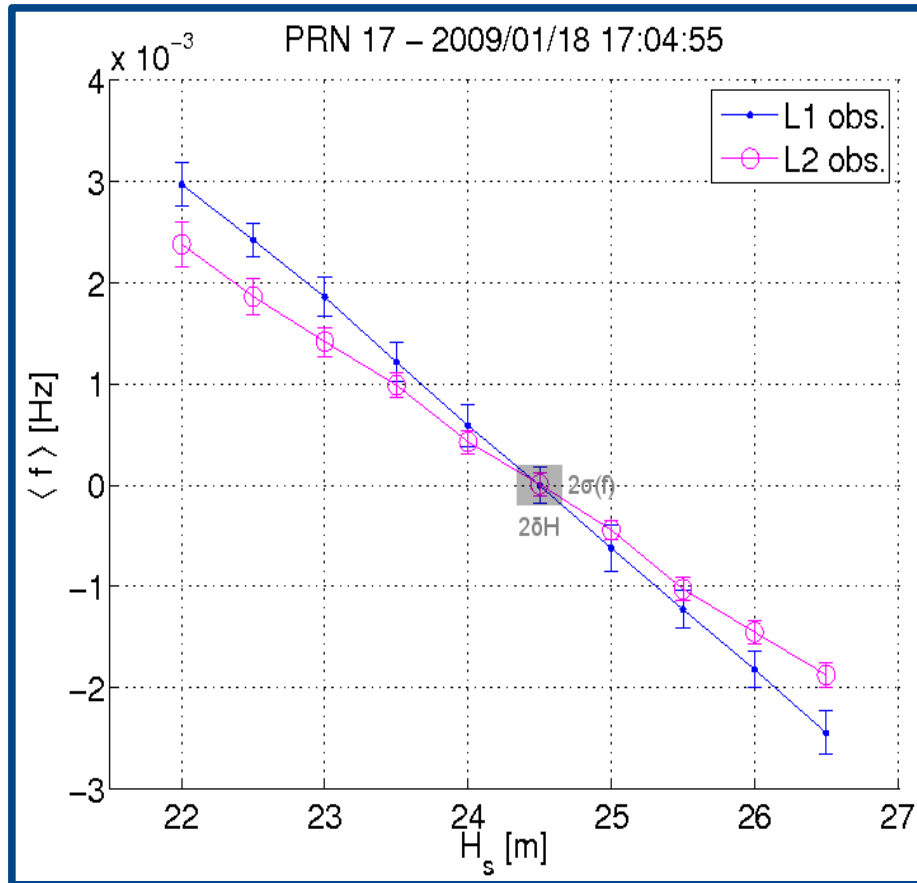


M. Semmling et al.,
"Detection of Arctic Ocean Tides using
GNSS-R"
in prep.

Altimetry

- residual phase method
- specular ocean heights
- several days
- semi-diurnal tides
- restriction due to refraction

Geometry



Ionospheric Delay

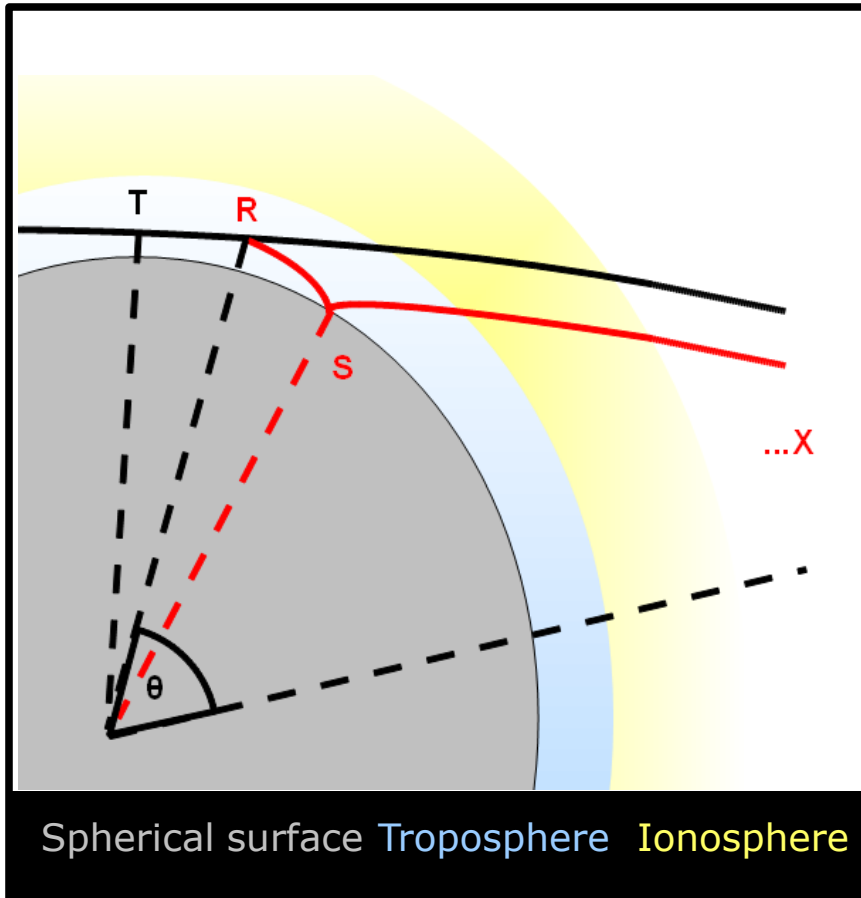
- dispersive refraction
- highly variable density of free electrons

Tropospheric Delay

- non dispersive refraction
- N_2 , O_2 , CO_2 const.
- highly variable H_2O content

Ray Tracing of Signal Path

Geometry



Ionospheric Delay

- dispersive refraction
- highly variable density of free electrons

Tropospheric Delay

- non dispersive refraction
- N_2 , O_2 , CO_2 const.
- highly variable H_2O content

Ray Tracing of Signal Path

Troposphere

$$N = k_1 \frac{p - e}{T} + k_2 \frac{e}{T} + k_3 \frac{e}{T^2}$$
$$N_w = k_2 \frac{e}{T} + k_3 \frac{e}{T^2}$$

$$k_1 = 0.7760 \frac{\text{K}}{\text{Pa}}$$

$$k_2 = 0.704 \frac{\text{K}}{\text{Pa}}$$

$$k_3 = 3.739 \cdot 10^3 \frac{\text{K}^2}{\text{Pa}}$$

M. Bevis et al., 1994
“GPS meteorology: Mapping zenith delays
onto precipitable water.”

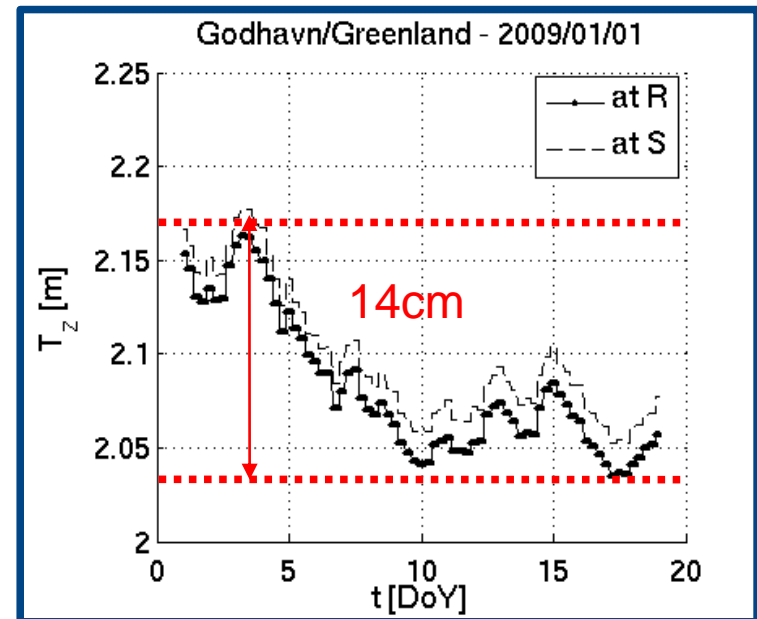
Neutral Atmosphere

- total refractivity N
- wet refractivity N_w

$$T_z = \int_R^X N(z) 10^{-6} dz$$

Total Zenith Delay

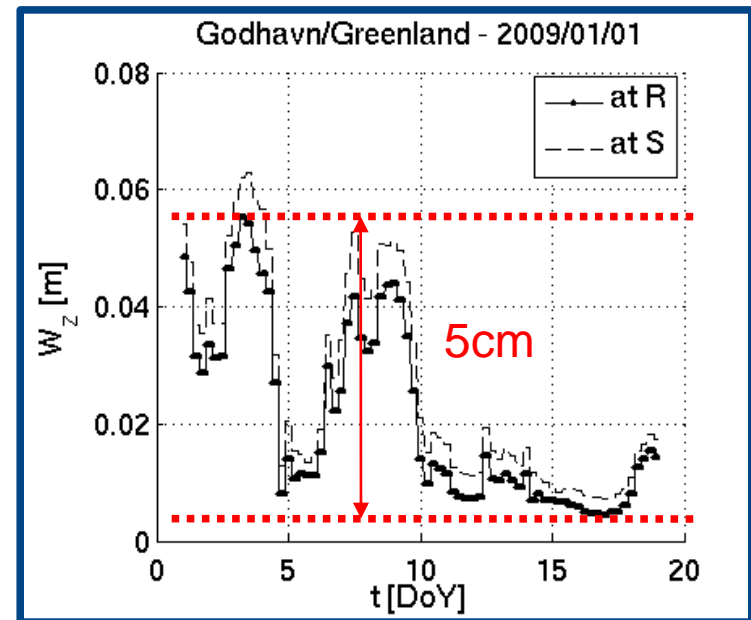
- prediction ECMWF
- at receiver
- at surface



$$W_z = \int_R^X N_w(z) 10^{-6} dz$$

Wet Zenith Delay

- prediction ECMWF
- at receiver
- at surface



Observation

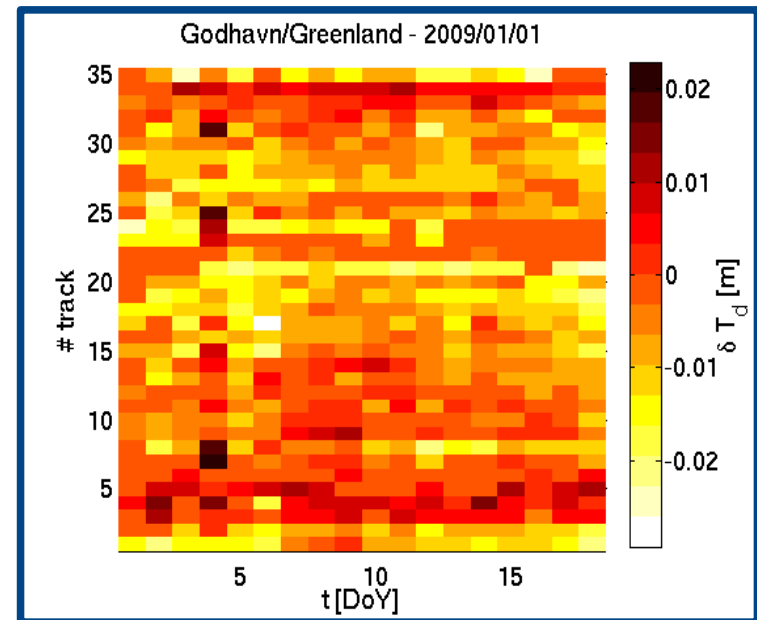
Slant Delay

- direct signal
- model T_d
- observation T

$$T_d = L_d - |XR|$$
$$\delta T_d = T_d - T$$

Error

- nearly anisotrop
- almost symmetric
- $\delta T/T < 1\%$



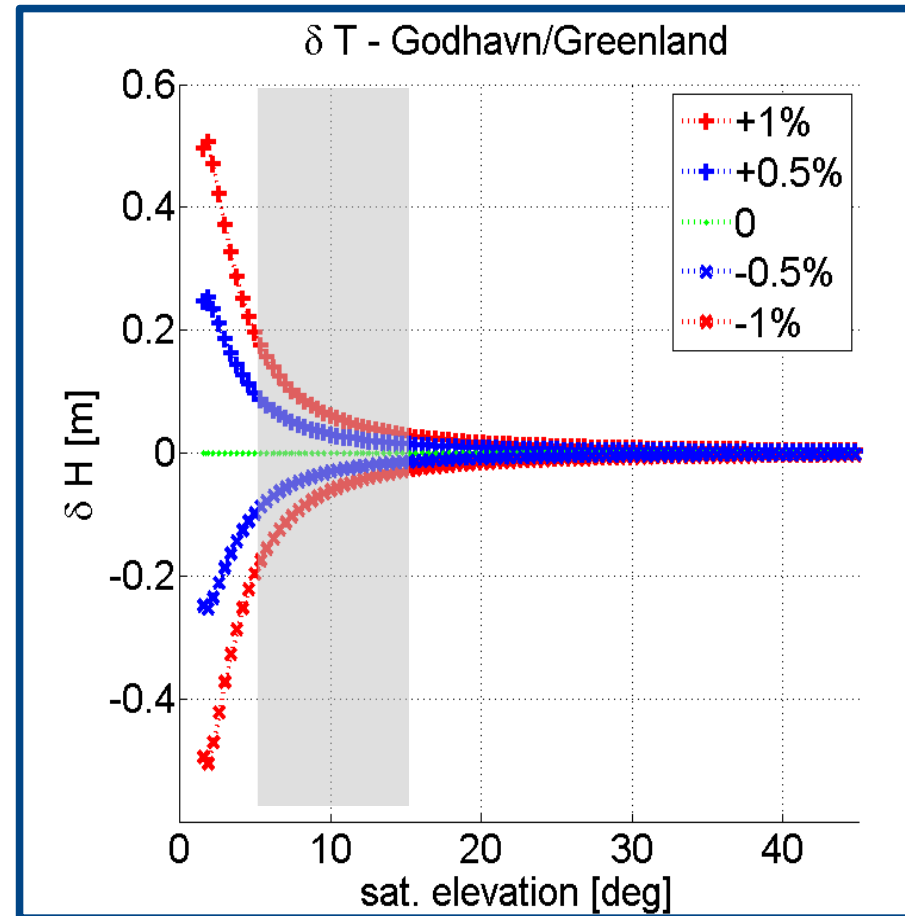
M. Bender et al., 2009
"Estimates of the information provided by
GPS slant data observed in Germany."

Results

$$\begin{aligned} f(H, T) &= 0 \\ \tilde{f}(\tilde{H}, T + \delta T) &= 0 \\ \delta H &= \tilde{H} - H \end{aligned}$$

Altimetric Bias

- over estimate
- under estimate
- symmetric
- <20cm

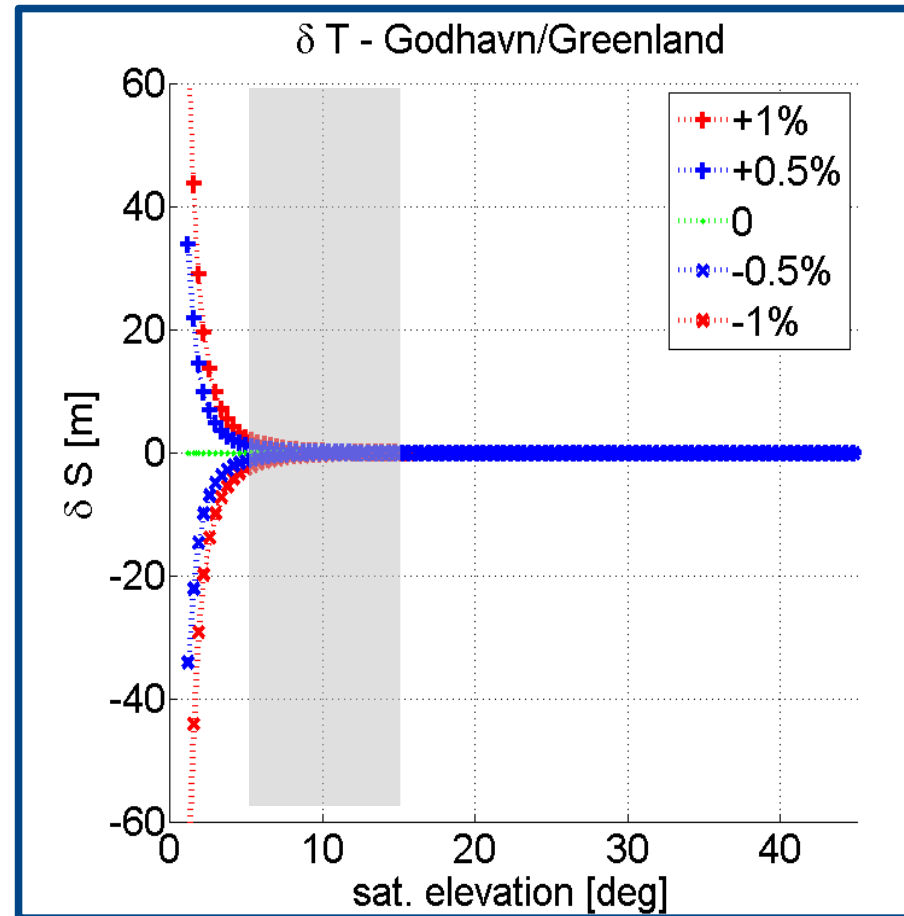


Results

$$\begin{aligned} f(s, T) &= 0 \\ \tilde{f}(\tilde{s}, T + \delta T) &= 0 \\ \delta S &= \overline{\tilde{S}S_t} \end{aligned}$$

Tangent Shift

- over estimate
- under estimate
- symmetric
- ~1m



Ionosphere

Model

$$\eta(z) = \eta_0 \exp \left[\frac{1}{2} \left(1 - \frac{z - z_0}{\Delta z} \right) - e^{-\frac{z - z_0}{\Delta z}} \right]$$

$$\eta_0 = 3 \text{ E}^{12} \text{ m}^{-3}$$

$$z_0 = 300 \text{ km}$$

$$\Delta z = 60 \text{ km}$$

F-Layer

- electron density η
- peak density η_0
- peak height z_0
- peak width Δz

S. Chapman, 1931

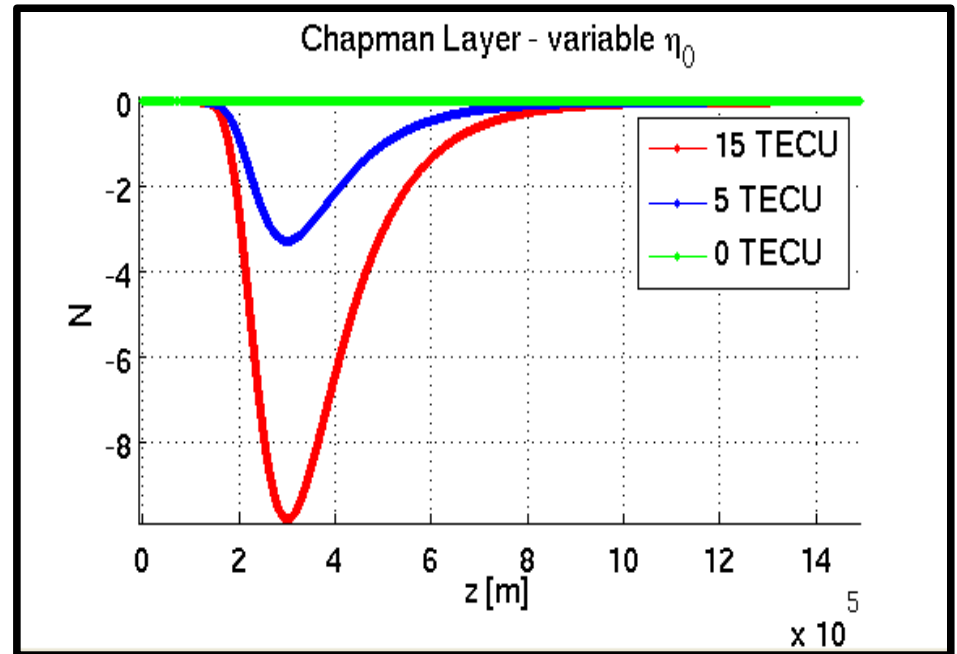
“Absorption and dissociative or ionising effects of monochromatic radiation in an atmosphere on a rotating earth.”

$$N = -\frac{A}{f^2} \eta$$

$$TECV = \int_R^X \eta(z) dz$$

Total Electron Content

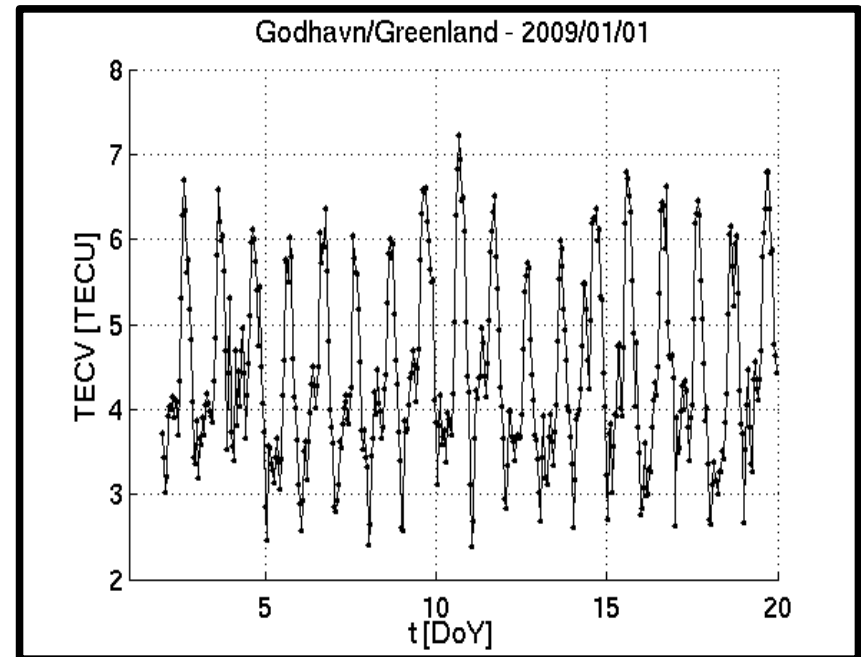
- vertical TEC
- scale η_0



Observation

Vertical TEC

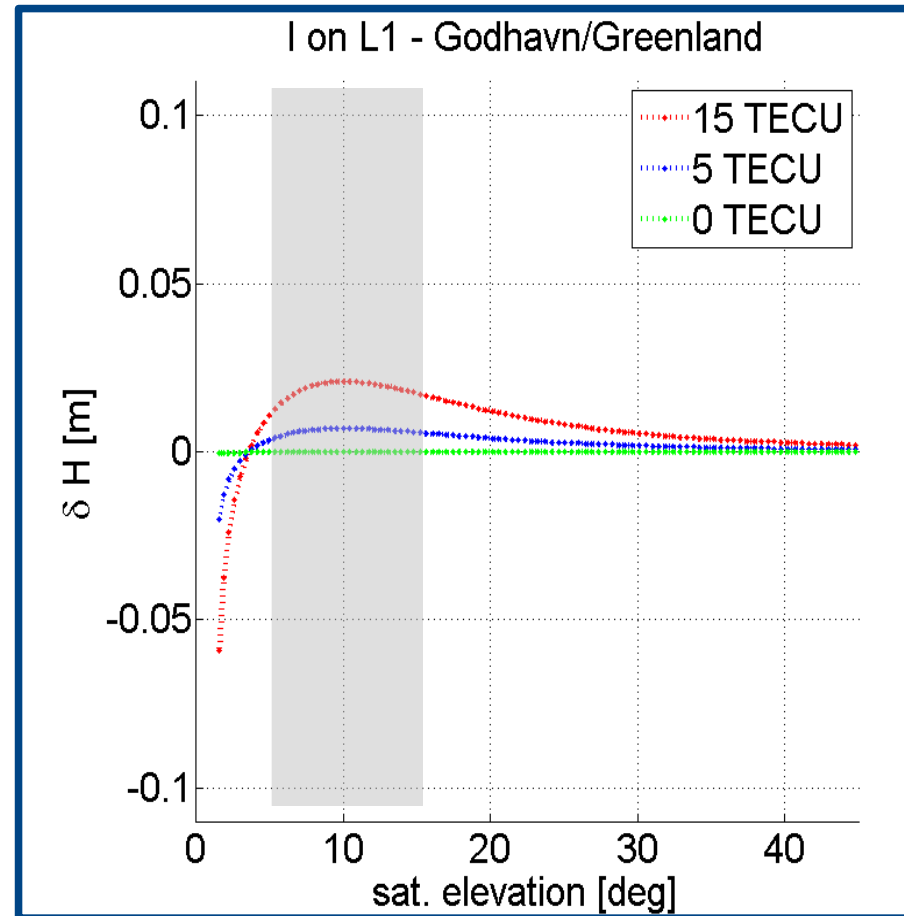
- direct signal
- mapped



$$\begin{aligned} f(H) &= 0 \\ \tilde{f}(\tilde{H}, I_{L1}) &= 0 \\ \delta H &= \tilde{H} - H \end{aligned}$$

Altimetric Bias

- different TECV
- slant elevations
- L1
- ~1cm

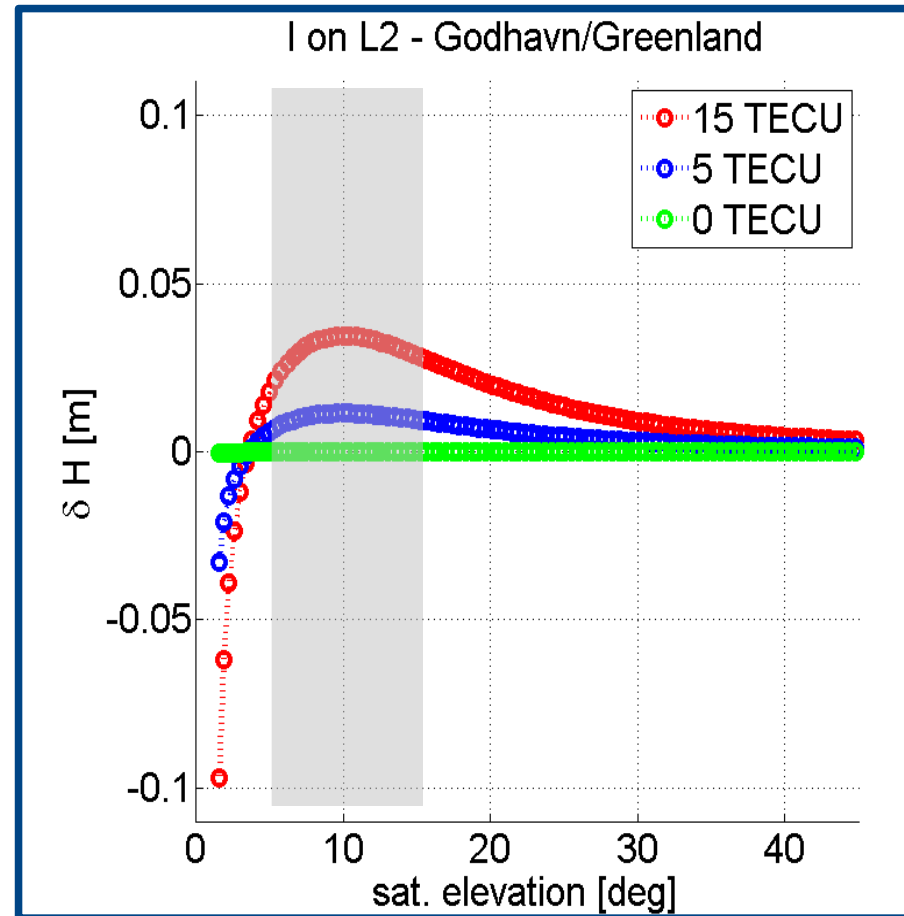


Results

$$\begin{aligned} f(H) &= 0 \\ \tilde{f}(\tilde{H}, I_{L1}) &= 0 \\ \delta H &= \tilde{H} - H \end{aligned}$$

Altimetric Bias

- different TECV
- slant elevations
- L2 larger
- ~1cm

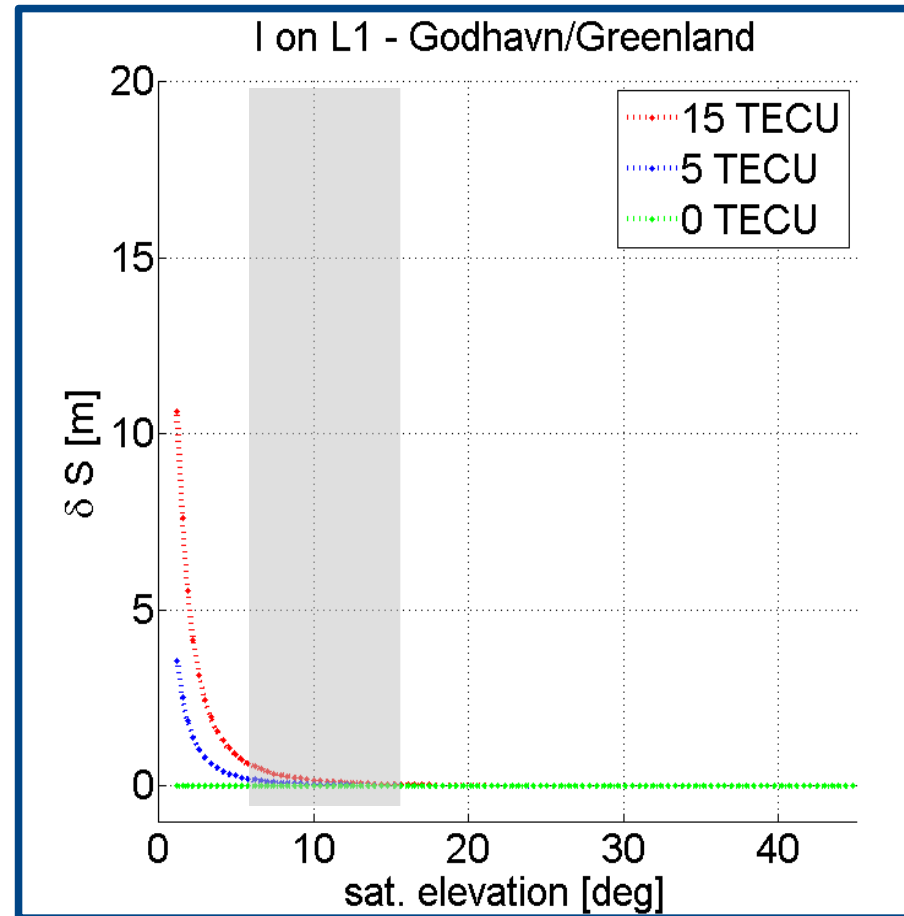


Results

$$\begin{aligned} f(s) &= 0 \\ \tilde{f}(\tilde{s}, I_{L1}) &= 0 \\ \delta S &= \overline{\tilde{s} S_t} \end{aligned}$$

Tangent Shift

- different TECV
- grazing elevations
- L1
- <1m

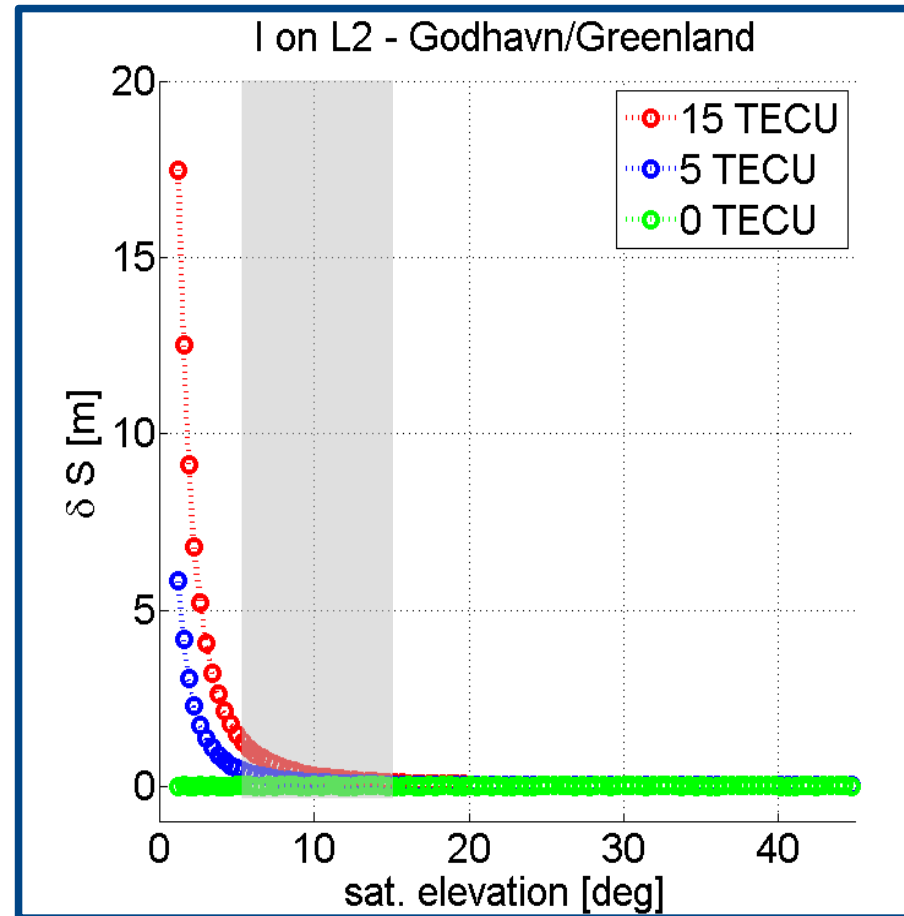


Results

$$\begin{aligned} f(s) &= 0 \\ \tilde{f}(\tilde{s}, I_{L2}) &= 0 \\ \delta S &= \overline{\tilde{s} S_t} \end{aligned}$$

Tangent Shift

- different TECV
- grazing elevations
- L2 larger
- <2m



Summary

Groundbased Experiment

- specular ocean height
- in remote sensing
- restricted by refraction

Airborne Proof of Concept

- less refraction
- few observations
- changes specular aircraft height

Next step to PaRIS

- new flight campaign
- longer observations
- measure geoid undulations

Outlook

Thanks to all colleagues.

Thank you, for your attention!

